

Determining ground layer turbulence statistics using a SLODAR-type method

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Joint work with:
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- SLODAR-type methods aim to recover the **vertical turbulence profile** from the cross-correlation of wavefront slope measurements from two guide stars
- The turbulence profile is crucial for AO systems with multiple guidestars, where strong prior information is required to stabilize the tomography problem
- SLODAR methods rely on the Kolmogorov/von Kármán models for turbulence statistics
- However, turbulence statistics at the ground can deviate from this model
- This issue is emphasised by the fact that often much of the turbulence strength is located close to the ground
- What if we could use the same measurements to both reconstruct the turbulence profile and infer a model for the ground layer turbulence simultaneously?

Introduction

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- SLODAR is a well-posed problem (fewer unknowns than measurements)
- Using inverse problems methods, we can extract more information from the same measurements
- The goal of our method is to simultaneously reconstruct the **vertical turbulence profile** of the atmosphere and the **power spectral density** at the ground layer
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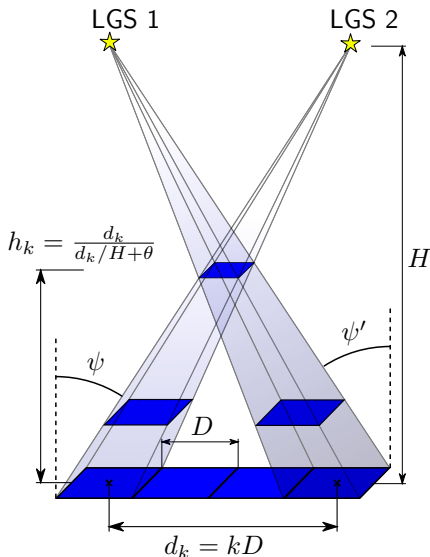
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Outline of this talk

- 1 Brief description of SLODAR
- 2 Solving for PSD with SLODAR
- 3 Numerical results
- 4 Open questions

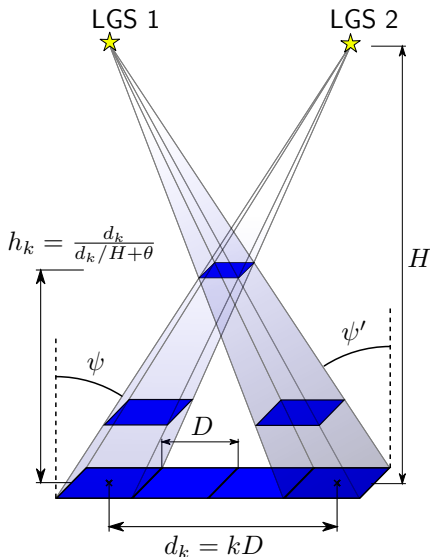
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- Measurement setup with two LGS wavefront sensors
- Correlation of measurements from two subapertures (one from each WFS) gives information about turbulence at an altitude depending on the distance between the subapertures
- SLODAR-type methods aim to reconstruct the vertical turbulence profile from these correlations
- Layer altitudes depend on LGS altitude H , LGS separation θ , and subaperture distances d_k



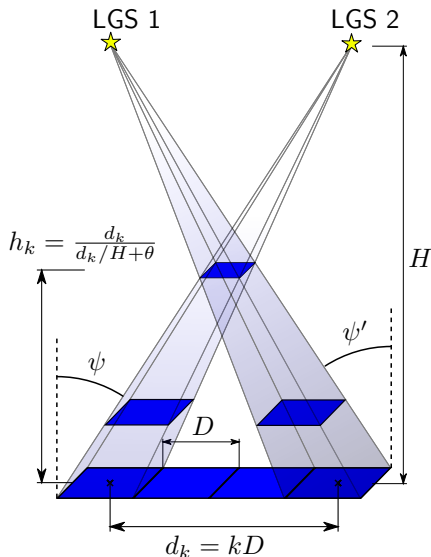
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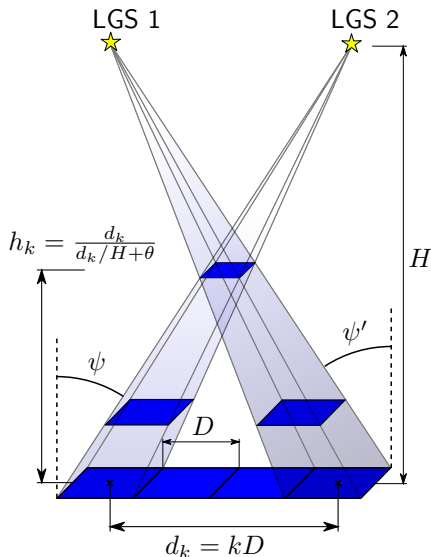
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Crash course to SLODAR

- Correlation of x -directional slope derivatives for subaperture separation \mathbf{d}_j :

$$\text{Cor}^x(\mathbf{d}_j) = \sum_k \frac{1}{D^4} \left(\int_{\mathbb{R}^2} e^{-2\pi i \boldsymbol{\xi} \cdot (\eta_k \mathbf{d}_j - h_k \boldsymbol{\theta})} \Psi_0(\boldsymbol{\xi}) |g_k^x(\boldsymbol{\xi})|^2 d\boldsymbol{\xi} \right) \rho_k$$

- ρ_k : turbulence strength at layer k
- $\Psi_0(\boldsymbol{\xi})$: von Kármán power spectral density, $\Psi_0(\boldsymbol{\xi}) = 0.0229(|\boldsymbol{\xi}|^2 + L_0^{-2})^{-11/6}$
- Collecting the integrals into a matrix \mathbf{A}^x (and similarly for y -slopes), we get

$$\mathbf{b} := \begin{pmatrix} \mathbf{b}^x \\ \mathbf{b}^y \end{pmatrix} = \begin{pmatrix} \mathbf{A}^x \\ \mathbf{A}^y \end{pmatrix} \boldsymbol{\rho} =: \mathbf{A} \boldsymbol{\rho}$$

- $\boldsymbol{\rho}$: vector of turbulence strengths ρ_k at altitudes h_k
- \mathbf{b} : correlations of WFS measurements with different subaperture distances \mathbf{d}_j
- Easy to solve in the least squares sense (e.g. using Matlab's `quadprog`):

$$\min_{\boldsymbol{\rho} \geq 0} \|\mathbf{A} \boldsymbol{\rho} - \mathbf{b}\|_2^2$$

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Modified SLODAR method

- Simple idea: for the ground layer $h_0 = 0$, replace Ψ_0 by an unknown PSD Ψ :

$$\rho_0 \Psi(\xi) = \sum_{l=0}^{N_r-1} \psi_l f_l(|\xi|)$$

- Note: ρ_0 included in above definition to avoid nonlinearity
- It follows that

$$A_{j0}^x \rho_0 = \frac{1}{D^4} \sum_{l=0}^{N_r-1} \left(\int_{\mathbb{R}^2} e^{-2\pi i \xi \cdot \eta_0 \mathbf{d}_j} |g_k^x(\xi)|^2 f_l(|\xi|) d\xi \right) \psi_l$$

- As before, collect the integrals into a matrix B to obtain

$$\mathbf{b} = \begin{pmatrix} \mathbf{b}^x \\ \mathbf{b}^y \end{pmatrix} = \begin{pmatrix} \mathbf{A}^x \\ \mathbf{A}^y \end{pmatrix} \boldsymbol{\rho} + \begin{pmatrix} \mathbf{B}^x \\ \mathbf{B}^y \end{pmatrix} \boldsymbol{\psi} =: \mathbf{A}\boldsymbol{\rho} + \mathbf{B}\boldsymbol{\psi}$$

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Modified SLODAR method

- Discretizing the PSD leads to many more unknowns than measurements
- The system $\mathbf{b} = \mathbf{A}\boldsymbol{\rho} + \mathbf{B}\boldsymbol{\psi}$ is strongly underdetermined \Rightarrow **inverse problem**
- Standard methods (e.g. least-squares solution) fail: noise in the measurement \mathbf{b} dominates
- Even worse, $\boldsymbol{\psi}$ decays rapidly since it is (close to) a power law
- Solution: **regularization!**
- Use **prior information** to favor solutions resembling a power law

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- **Method 1: Parametric Power Law.** Seek solutions of the form

$$\psi_j(c, \gamma) = c(r_j^2 + 1/L_0^2)^{-\gamma}$$

- Least-squares solution:

$$\min_{\rho, c, \gamma \geq 0} \|A\rho + B\psi(c, \gamma) - \mathbf{b}\|_2^2$$

- Regularization by discretization
- Nonlinear, but low number of parameters makes it easy to solve

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- **Method 2: Tikhonov Regularization.** Solve least-squares problem

$$\min_{\rho \geq 0, \psi \geq 0} \|A\rho + B\psi - \mathbf{b}\|_2^2 + \beta \|\mathbf{\Gamma}(\psi - \psi_0)\|_2^2 + \beta' \|\mathbf{\Gamma}'(\psi - \psi_0)\|_2^2$$

- $\mathbf{\Gamma}$, $\mathbf{\Gamma}'$ and ψ_0 encode our prior information about the unknown PSD ψ
- β and β' control how strongly prior information is enforced
- Prior PSD ψ_0 given by method 1
- Intuition for choosing $\mathbf{\Gamma}$ and $\mathbf{\Gamma}'$: impose a penalty on the relative error between ψ and ψ_0 , and similarly between their derivatives
- Easy to solve using quadratic programming
- Slightly more complex than method 1, but much more suitable for cases where true solution is not quite a power law

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Numerical simulations

- Simulations using MOST, an AO system simulation toolkit developed at JKU
- Telescope diameter 42m
- WFSs with 84×84 subapertures (i.e. $D = 0.5\text{m}$)
- LGS separation $\theta = (7.5, 0)$, in arcminutes
- $N_L = 61$ layers
- $d_k = (k, 0)$ with $k = 0, \dots, 60$
- $H = 90\text{km}$
- $h_k = \frac{kD}{kD/H + |\theta|} \Rightarrow h_0 = 0\text{m}, h_{60} \approx 12\text{km}$
- Measurements over 50 timesteps
- Ground layer PSD discretized by 400 points r_i , with $r_{\max} = 10$

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- $h_k = \frac{kD}{kD/H + |\theta|} \Rightarrow h_0 = 0\text{m}, h_{60} \approx 12\text{km}$
- Measurements over 50 timesteps
- Ground layer PSD discretized by 400 points r_i , with $r_{\max} = 10$

- Atmosphere simulated as discrete layers at the altitudes h_k used by SLODAR
- von Kármán power spectral density given by

$$\Psi_0(\xi) = 0.0229(|\xi|^2 + 1/L_0^2)^{-11/6}$$

- For ground layer $k = 0$, we changed the PSD to

$$\Psi(\xi) = 0.0229(|\xi|^2 + 1/L_0^2)^{-1.5732}$$

- To test method 2, we also simulated data where three small "bumps" were added to the above PSD

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PSD reconstruction with method 1

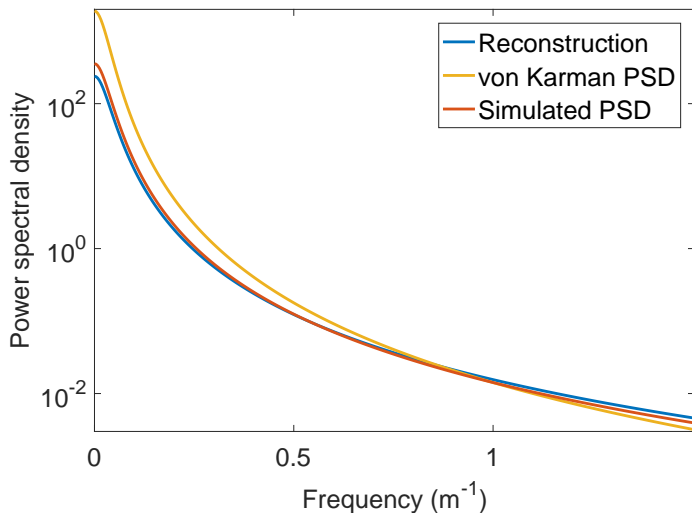
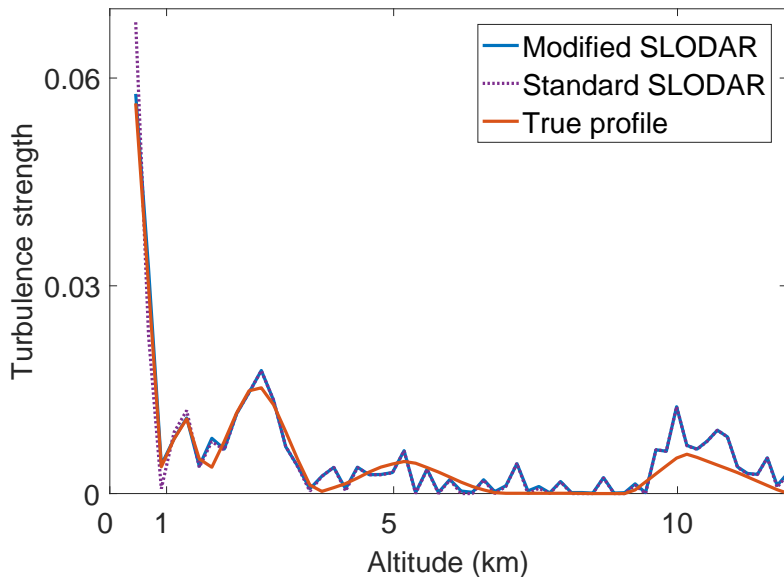
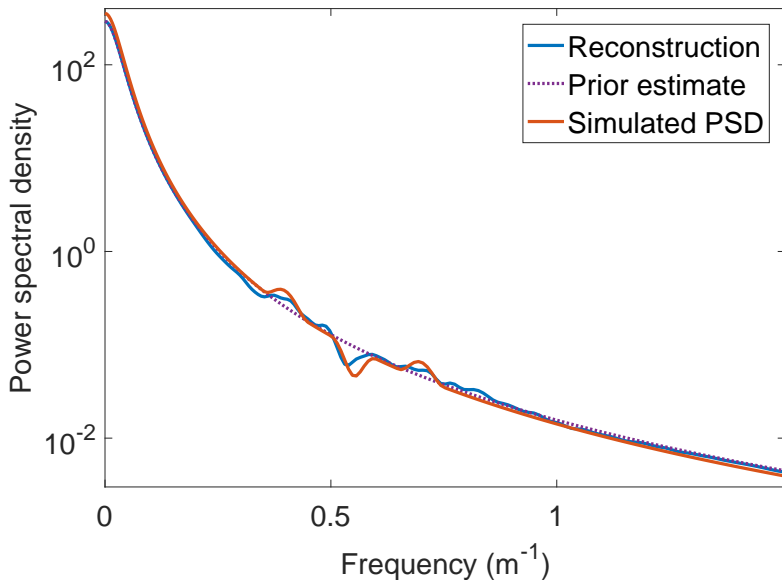


Figure: Simulated PSD exponent -1.5732, reconstructed exponent -1.497.

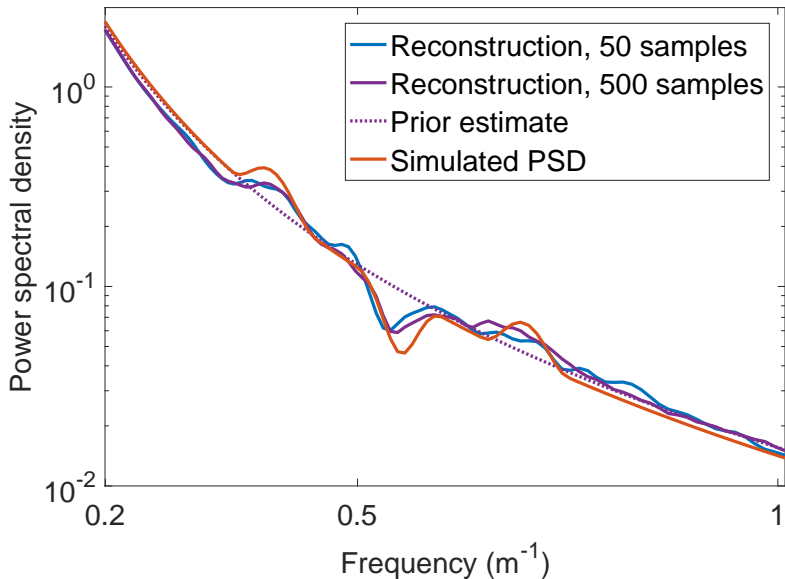
Turbulence profile reconstruction with method 1



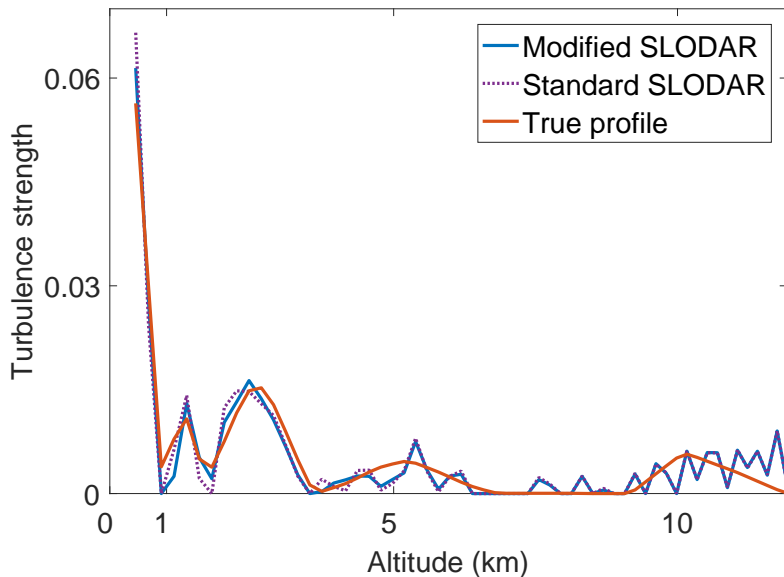
PSD reconstruction with method 2, using 50 timesteps



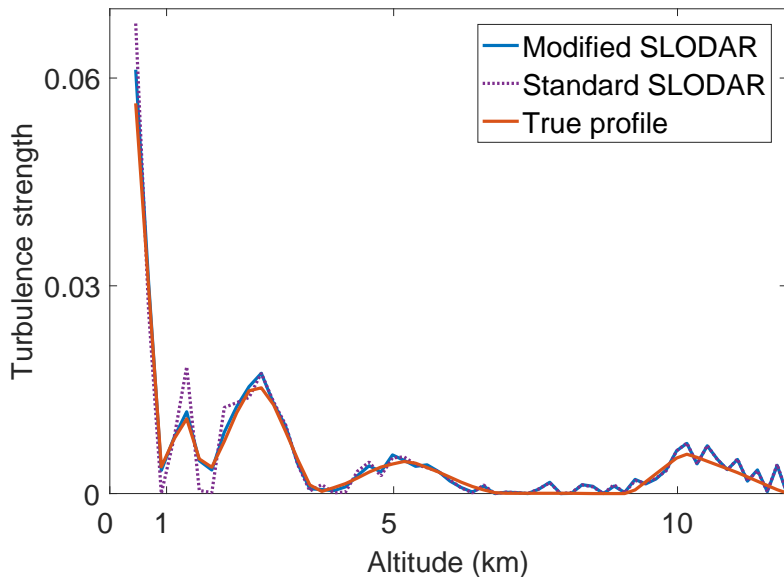
PSD reconstruction with method 2, using 500 timesteps



Turbulence profile, method 2, using 50 timesteps



Turbulence profile, method 2, using 500 timesteps



- What is a realistic class of models for the ground layer turbulence statistics?
- How much does real low altitude turbulence deviate from an isotropic and homogeneous model?
- How much data do we need for a good reconstruction? How quickly/slowly do the turbulence statistics for the bottom layer change? In other words, do we have enough time to gather the necessary amount of data?
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Thank you for your attention!